1	Dynamics of low and high density microplastics in the world's largest choked coastal				
2	lagoon under contrasting meteoceanographic conditions				
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12 ABSTRACT

13 The contamination by plastic waste in aquatic environments has become a global issue, 14 scientifically reported since 1970. The size and durability of microplastics (MPs, $1 \mu m >$ 5 mm) have made these debris widely distributed in aquatic environments. Despite 15 16 various ongoing initiatives, there is a need to fill gaps in understanding how MPs are transported from their release sources to their final destinations. Therefore, understanding 17 the distribution and dynamics of MPs in coastal areas, such as lagoons and estuaries, 18 which are considered continental sources of MPs to the oceans, is essential to help fill 19 these gaps and propose alternatives for managing what is the environmental problem of 20 21 the century. In this context, this study aimed to assess the transport patterns of MPs in 22 Patos Lagoon, the largest choked coastal lagoon in the world, considering contrasting 23 meteoceanographic conditions in the system dynamics, such as wind and discharge. Using the hydrodynamic model TELEMAC-3D and the model for plastics particles 24 25 TrackMPD, simulations were performed using a type of MP polymer (Polypropylene -26 PP). The simulations of MP transport considered advection, dispersion and the 27 contribution of biofilm in increasing particle density (representing high-density microplastics). The results indicated a gradient of MPs retention from north to south, with 28 higher concentrations of MPs occurring in the northern part of the system. The central 29 30 region of the lagoon showed a greater tendency for MP export towards the southern region than retention. Meanwhile, the estuary region of Patos Lagoon exhibited a tendency for 31

export of low-density MPs free of biofilm and retention of higher-density MPs, subject to the action of the Plastisphere. Additionally, based on the results obtained from density occurrence maps, it was possible to suggest potential accumulation areas of MPs throughout the lagoon system, reinforcing that the system can act as a sink in specific regions.

Keywords: plastic contamination; plastic accumulation; hot spots; estuaries; coastallagoons; numerical modelling.

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40 1. INTRODUCTION

The presence of plastic debris in aquatic environments has been documented in the 41 scientific literature since the early 1970s, and in contemporary times, plastic 42 43 contamination became ubiquity in these environments. The origin of plastics in the oceans can be attributed to both continental and maritime sources. According to Andrady (2011), 44 45 approximately 80% of this contribution is linked to microplastics from continental sources (Li et al., 2018), which are highly susceptible to enter the aquatic environment 46 47 through rivers, surface runoff, and wastewater systems because of their low density and weight (Horton et al. 2017). Meijer et al. (2021) estimated that between 0.08 and 2.7 48 million tons of plastics enter the oceans annually through rivers. 49

Among the different aspects related to the problem, special attention can be given to 50 microplastics (MPs), which encompasses a diverse array of microparticles and fibers, 51 typically ranging in size from 1 µm to 5 mm (Hanvey et al., 2017). Unlike their larger 52 plastic counterparts, MPs may be of primary origin, intentionally manufactured in this 53 54 size range (as pellets), or of secondary origin, resulting from the gradual breakdown of larger plastic items within the environment (Horton et al., 2017). Their small size and 55 56 enduring nature present distinct challenges, as they can easily infiltrate and persist within marine and coastal ecosystems indefinitely. From surface waters to the ocean floor, 57 58 microplastics have been documented (Barnes et al., 2009; Ling et al., 2017, Pan et al., 2019; Abel et al., 2021; Erkan et al., 2023; Tsuchiya et al., 2024), directly and indirectly 59 affecting these environments. Their presence give rise to detrimental consequences, 60 including alterations in habitats (Gall & Thompson, 2015), disruptions to socio-economic 61 62 activities, adverse effects on aquatic organisms (Wright et al., 2013; Koelmans et al., 63 2016), and even potential implications for human health (Napper et al., 2019).

Despite various initiatives to understand the transport and distribution patterns of MPs in 64 65 different aquatic compartments, obtaining a comprehensive global view of their distribution remains challenging. Discrepancies between the estimated amount of plastic 66 floating on the ocean surface (Eriksen et al., 2014) and the estimated amount entering the 67 ocean (Jambeck et al., 2015) suggest that some of this debris may be retained near its 68 source regions. With approximately half of the world's population living within 80 km of 69 the coast, coastal areas are already indicated as sources of plastics to the oceans (Cole et 70 71 al., 2011) as well as accumulation environments (sinks). High concentrations have been 72 observed in port areas (Auta et al., 2017), as well as salt marshes (Weinstein et al., 2016; 73 Pinheiro et al., 2021) and estuaries (Díez-Minguito et al., 2020; Forero-López et al., 2021; 74 Shabaka et al., 2022; Banik et al., 2024).

75 The distribution of these debris can be complex, with a significant portion remaining 76 suspended in the water column (Ikenoue et al., 2023), while some may sink (Lobelle et al., 2021), and others may be transported to beaches (Ranjani et al., 2022) and coastal 77 78 areas. Generally, microplastics with a density range lower (higher) than the surrounding fluid are more likely to float (sink) in the water column (Waldschläger & Schüttrumpf, 79 2019), being considered low-density (high-density) MPs. Low-density MPs, such as 80 polypropylene, have a density range of 0.90-0.99 g/cm³; however, particle density may 81 vary due to biofouling development (Chubarenko et al., 2016), resulting in high-density 82 83 MPs. Among the processes that can be responsible for the removal of MPs from the surface, one of the most relevant is biofouling (Kaiser et al., 2017). This can be defined 84 as the colonization of micro to macro-organisms on the surface of a consolidated 85 substrate, and it is called Plastisphere when formed on plastic debris (Zettler et al., 2013, 86 87 Su et al., 2022).

The monitoring of MPs distribution in coastal ecosystems is crucial for obtaining insights 88 into the transport and accumulation patterns of these particles, which are of utmost 89 importance for the proper management of these pollutants (Kershaw et al., 2015). 90 However, monitoring studies are conducted through sampling efforts, which involve high 91 financial costs and time demands, especially in larger areas. Due to the challenges 92 93 associated with monitoring this type of solid waste, the numerical modeling approach has 94 emerged as a promising and alternative tool for understanding the distribution of MPs in extensive areas and in the water column (Isobe et al., 2009; Lebreton et al., 2012; Critchell 95 96 & Lambrechts, 2016; Jálon-Rojas et al., 2019; Díez-Minguito et al., 2020; Baudena et al.,

97 2022). This approach exposes these particles to different physical processes, allowing for
98 tracking and providing insights into their origins, distribution and destinations. Moreover,
99 it contributes to more accurate mapping of potential risk areas and suggests monitoring
100 sites. Numerical models also facilitate the evaluation of hypothetical scenarios and the
101 consistent interpretation of sparse data.

102 Patos Lagoon is considered the largest choked coastal lagoon in the world (Kjerfve, 103 1986), serving as a central hub for population, commerce, industry and recreation. As a 104 result, it is also subject to the disposal of industrial and municipal waste. According to 105 Santos et al., (2023), between the years 2010 and 2017, the annual input of plastic waste 106 into Patos Lagoon was estimated between 21.67 and 107.19 thousand tonnes, with the 107 main polymers found being Polypropylene (PP), Polyethylene (PE), and Polyvinyl 108 Chloride (PVC). In the estuarine area is located the second largest harbor of Brazil, Rio 109 Grande Harbor. This industrial area is responsible for the transport and operation of up to 50 million tons per year, subjecting this region to large amounts of urban waste (Pinheiro 110 111 et al., 2021).

112 Recent studies have highlighted the occurrence of plastic debris in different areas of Patos Lagoon (Fig. 1) (Silva & Sousa, 2021; Alves et al., 2022; Pinheiro et al., 2022; Santos et 113 114 al., 2023). Rodriguez et al. (2024) simulated the export of MPs through the Patos Lagoon 115 coastal plume and their accumulation tendency in the coastal zone, but despite these 116 efforts, studies focusing on the transport of MPs in the lagoon main water body remain limited due to its length (240 km). In this context, the aim of this study was to investigate 117 118 the advection and dispersion transport process of both low and high-density MPs in Patos Lagoon, considering contrasting meteoceanographic conditions and high-density MPs 119 120 resulting from biofouling formation and its implications for the transport patterns. This study was carried out with the TrackMPD model. 121

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123 2. MATERIAL AND METHODS

124 **2.1 STUDY AREA**

Patos Lagoon is a shallow coastal plain choked lagoon located in southern Brazil (30° –
32° S, 50° – 52° W) (Fig. 1). With an area of approximately 10.230 km², extending over
240 km, and with an average depth of 5 m, the lagoon drains a basin of 200.00 km².
Predominantly consisting of freshwater, the main body of the lagoon experiences coastal

water influence within the initial 60 km from the lagoon's mouth, defining the Patos
Lagoon Estuary (Fig. 1C). The estuarine area is the deepest region of the system due to
the presence of natural and artificial channels (Möller et al., 2001), and it is also the only
connection to the Atlantic Ocean through a long (20 km) and narrow (500 m) channel
fixed by a pair of 4 km long jetties (Fig. 1C) (Franzen et al., 2023).

The lagoon hydrodynamics is driven by continental discharges and the action of both 134 135 remote and local winds (Moller et al., 2001; Vaz et al., 2006). The main tributaries of the lagoon are Guaíba and Camaquã rivers and São Gonçalo Channel. The continental 136 137 discharges in the lagoon have seasonal variability typical of mid-latitudes, with high discharges in austral winter and late spring and low discharges during austral summer and 138 139 fall and with interannual (Távora et al., 2020) and interdecadal variations (Bortolin et al., 2022). The lagoon is also under the influence of ENSO cycles, resulting in higher (El 140 141 Niño) or lower (La Niña) precipitation and continental discharges (Távora et al., 2019).

Although the lagoon circulation is influenced by both local and remote wind effects, the 142 local winds act as the primary driving force of the main lagoon body (Fig. 1B) when 143 discharges are < 2000 m³.s⁻¹, producing a set-up/set-down mechanism. Meanwhile, non-144 local winds tend to promote the exchanges between the estuarine area and the adjacent 145 146 continental shelf (Fig.1C) due to the Ekman Transport acting 90° of the wind direction. 147 Throughout the year, northeasterly (NE) winds dominate, while southwesterly (SW) become more important during autumn and winter, as frontal systems become more 148 frequent (Möller et al., 2001). This regional wind pattern coincides with the main axis of 149 150 the lagoon (NE - SW). Moller et al. (1996) indicated that easterly (E) winds occurring during summer and spring are a result of sea breezes. 151



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Figure 1: Study area (A, B, C). Model domain with the main tributaries of the lagoon
(outline dots), bathymetric information represented in the color scale, identification of
open boundaries, and applied boundary conditions for the simulations (B). Patos Lagoon
estuary (C).

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158 2.2 HYDRODYNAMIC MODEL

159 The TELEMAC-3D hydrodynamic model. version V7P0, was applied (http://www.opentelemac.org) to generate the advective field for the MPs dispersion 160 161 simulations with TrackMPD. TELEMAC-3D was developed by the Laboratoire National d'Hydraulique (LNH) of Electricite de France (EDF). Based on the Finite Element 162 163 Method, it solves the Reynolds averaged three-dimensional Navier-Stokes equations, 164 considering local variations in the free surface of the fluid over time and hydrostatic 165 pressure, in addition to Boussinesq approximations to solve the motion equations 166 (Hervouet, 2007). The resulting equations are suitable for application in shallow waters, where vertical fluid velocities are much smaller than horizontal velocities. Advection and 167 diffusion equations are used for calculating the transport of tracers such as temperature, 168 salinity and suspended sediment. The key outcomes obtained at each point of the 169 170 numerical grid include velocities in the three directions (u, v, and w), concentrations of 171 transported tracers, and the elevation of the free surface of the fluid.

The numerical domain of the model encompasses the entire lagoon and extends to a depth of 2.300 m in the adjacent coastal region (Fig. 1B), being discretized through a non-

structured grid of finite elements that allows more control of the element distribution 174 175 within the domain, and consequently, more autonomy to establish finer resolution in regions with complex morphology or bathymetry. Vertical discretization was achieved 176 177 through 7 sigma layers. The non-structured grid used in this study comprises 52.098 points and was generated through the BlueKenue software, version 3.3.4, 178 179 (https://nrc.canada.ca/en). The hydrodynamic performance of TELEMAC for the study area has been extensively calibrated and validated by various authors (Fernandes et al., 180 181 2002, 2005, 2007, 2021; Bitencourt et al., 2020, Lisboa et al., 2022, Franzen et al., 2023), 182 consistently showing good to excellent results for both RMSE and RMAE. The resulting 183 grid has already been calibrated and validated by Silva et al. (2002), who also presents 184 details about the configurations used for this simulations.

185 Initial and boundary conditions for the numerical grid were extracted from different 186 sources, interpolated, and prescribed at the respective open boundaries of the grid (Fig. 1B). Throughout the three-dimensional domain, initial fields of salinity and temperature 187 188 were prescribed based on the results of the Hybrid Coordinate Ocean Model (HYCOM, https://hycom.org) + NCODA, which have a temporal resolution of 1 day and a spatial 189 190 resolution of 1/12°. Initial concentrations of suspended sediment were considered null throughout the domain. The OSU Tidal Inversion System (OTIS) was used for 191 prescribing sea surface elevation at oceanic boundaries. Sea surface elevation is 192 193 calculated by the inverse solution of Laplace's equations for tidal dynamics, using data 194 collected by the TOPEX/POSEIDON Project, internally available in the TELEMAC-3D 195 model. To obtain a reliable estimate of sea surface elevation, 33 tidal harmonic 196 components were considered. Once the sea surface elevation is known at one edge point, 197 the Continuity Equation is integrated, and the components of current velocity are obtained. At the surface boundary, the numerical grid was forced with results from the 198 199 global ECMWF model (European Center for Medium-Range Weather Forecast, http://www.ecmwf.int), with a temporal resolution of 24 h and spatial resolution of 0.75°. 200 201 At the continental boundaries, daily river discharge data from the National Water Agency 202 (ANA, http://www.ana.gov.br) for Guaíba River and Camaquã River were used. For the 203 São Gonçalo Channel discharge, water level data obtained from the Mirim Lagoon Agency (ALM, https://wp.ufpel.edu.br/alm/agencia) were used and transformed into 204 205 daily discharge from a rating curve (Oliveira et al., 2015).

The period of interest for our study corresponds to the entire year of 2013 (365 days), saved every 24 hours. This specific year was chosen based on the Oceanic Niño Index (ONI) analysis, where it appears as a neutral year with respect to the warm and cold phases of the ENSO phenomenon. Therefore, selecting the year 2013 allows studying the MPs transport process in the Patos Lagoon and its estuary under hydrodynamic conditions without the influence of the ENSO phenomenon. To accomplish this, the hydrodynamic model results were used to force the TrackMPD model (Jalón-Rojas et al., 2019).

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214 2.3 LAGRAGIAN TRACKING MODEL

215 The Tracking Marine Plastic Debris model (TrackMPD, Jálon-Rojas et al., 2019), is a 216 particle transport model that accepts current velocities from various hydrodynamic models such as MOHID, POM, MARS, NEMO, and TELEMAC-3D. The hydrodynamic 217 velocities that the model supports can result from the discretization of structured or 218 219 unstructured grids. Regarding the vertical discretization of the domain, TrackMPD 220 accepts hybrid vertical coordinates $(z-\sigma)$ and sigma coordinates (σ) , depending on the 221 settings used in the hydrodynamic model. It takes into account the physical characteristics of each type of plastic, such as size, shape, particle density and settling velocity. 222 223 Additionally, processes influencing the behavior and transport of MPs, such as advection, 224 turbulence, deposition, beaching, resuspension, biofouling, and degradation, can be 225 independently added to the model, according to the purpose of the study.

TrackMPD can be used in both 2D and 3D modes. Its operation occurs in two stages, where the first involves the integration and storage of velocity fields derived from the hydrodynamic model, and the second consists of simulations involving the calculation of particle trajectories. Once the velocities have been converted to TrackMPD's standard input formats (step 1), different simulations can be conducted using various model configurations (step 2), without the need for new integrations of the hydrodynamic results.

As the model is based on a Lagrangian approach, each particle is advected using the velocity fields, where the latter is interpolated in both time and space for the particle's location. Advection can be calculated using the second or fourth-order Runge-Kutta method (RK2 and RK4). Higher orders in the method, despite exhibiting better precision, entail higher computational costs. Additionally, considering the temporal and spatial

scales of the studied physical problem, it is necessary to configure the integration time 238 239 interval for the calculation of particle advection. TrackMPD resolves calculations at two 240 integration timescales, an external and an internal one. The external timestep (Δt_e) is 241 derived from the hydrodynamic model and corresponds to its outputs (e.g., results every 242 24 h). The internal timestep (Δt_i) is used for calculating particle trajectories (e.g., 30 243 minutes). To compute advection accurately and avoid an exaggerated estimate of particle displacement, it is recommended to use a smaller internal timestep compared to the 244 245 external one.

246 The behavior of particles in the water column depends on their physical characteristics, which determine their settling velocity (w_s , m.s⁻¹). Settling velocity occurs when the 247 248 gravitational force, minus the particle's buoyancy, is greater (smaller) than the drag force. Thus, particles exhibit positive buoyancy (rise) or negative buoyancy (settling). The 249 250 settling velocity of MPs can be constant or vary over time due to the influence of degradation and biofouling formation, among other processes. Therefore, particle 251 trajectories in a three-dimensional domain are determined by advection, diffusion, and 252 253 their vertical displacements through the following equations:

254
$$dX(t) = dX_{adv}(t) + dX_{diff}(t) = U(x, y, z, t)dt + dX'(t)$$
(1)

255
$$dY(t) = dY_{adv}(t) + dY_{diff}(t) = V(x, y, z, t)dt + dY'(t) (2)$$

256
$$dZ(t) = dZ_{adv}(t) + dZ_{diff}(t) + dZ_{dep}(t) = W(x, y, z, t)dt + dZ'(t) - ws(t)dt (3)$$

where $dX_{adv} = (dX_{adv}, dY_{adv}, dZ_{adv})$ represents advection established in the zonal (x), meridional (y), and vertical (z) axes obtained through the velocity fields of the hydrodynamic model U = (U, V, W). Turbulent diffusion $dX_{diff} = (dX_{diff}, dY_{diff}, dZ_{diff})$ is represented by a random component dX' = (dX', dY', dZ') added to the particle's movement on a much smaller scale than the advection movement. Vertical displacements dZ_{dep} depend on the settling velocity (w_s) of MPs.

It is also possible to calculate the settling velocity through formulations provided in the model, which result from experiments involving MPs (Khatmullina & Isachenko, 2017; Waldschlager & Schuttrumpf, 2019; Jalón-Rojas et al, 2022). These experiments serve as a basis for more realistic parameterizations of the variables. Through these formulations, it is also possible to include the effect of biofouling on particles, which alters the settling velocity (w_s). By defining a biofilm growth rate as a function of time, it is possible to calculate the density rate varying over time (g.cm³.day), defined as:

$$Ps = P_0 + DR\Delta t_i$$

where P_0 is the polymer density, Δt_i is the internal timestep, responsible for calculating displacements, and *DR* is the biofilm growth rate as a function of time. A more detailed description of the model's equations and all available numerical processes and solutions is presented by Jálon-Rojas et al. (2019). The TrackMPD validation for Patos Lagoon is presented by Rodriguez et al. (2024). Simulations were carried out using TrackMPD version 2.2.

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278 2.4 DENSITY PROBABILITY MAPS

279 Using a statistical approach to better highlight the accumulation trends of MPs (hotspots), 280 results from the TrackMPD model were analyzed through probability density maps, calculated using probability density functions (PDFs). These density maps provide a 281 282 visual representation of potential regions of MP accumulation in the system, identifying areas with the highest probability of particle occurrence. According to Mitarai et al. 283 (2009), this type of analysis can be applied to determine expected tracer concentrations, 284 widely used for predicting dispersion patterns. The functions are calculated based on the 285 286 probability of a particle moving from one location to another during a time interval, counting the number of particles per interval and subsequently normalizing by the total 287 number of the particles, mapping these through the grouping of particle positions into 288 histograms (Jálon-Rojas et al., 2019). The PDFs were calculated at a grid resolution of 289 approximately 0.03 km². 290

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292 2.5 SIMULATIONS SETUP

The simulations were carried out considering high (scenarios 2 and 4) and low (scenarios 1 and 3) continental discharge in Patos Lagoon (Tab. 1). In Scenarios 1 (low discharge) and 2 (high discharge) the trajectories of low-density MPs were simulated considering the entire lagoon, with MPs particles being released from the three main tributaries of the system (Guaíba and Camaquã rivers and São Gonçalo Channel).

According to Santos et al. (2022), an estimate of the contribution of plastic waste from 298 299 the tributaries into Patos Lagoon suggests that Guaíba and Camaquã rivers, and São 300 Gonçalo Channel introduce percentage contributions of 60%, 20%, and 20% of the total 301 load of plastic waste into the system, respectively. Thus, for Scenario 1 and 2, covering 302 the entire lagoon, 2.000 particles were proportionally distributed among the three 303 tributaries (1.200, 400 and 400), and for Scenario 3 and 4, 400 particles were released from São Gonçalo Channel (Tab 1). The authors also reported that among all types of 304 plastics described, the most abundant type was polypropylene (PP), which corroborates 305 306 with the results described by Rodriguez et al. (submitted). Therefore, in all simulated 307 scenarios, the chosen polymer type was PP.

308 For scenarios 1 and 2 in the main lagoon, the only physical processes taken into account were advection and dispersion. For Scenarios 3 and 4, the transport and accumulation of 309 310 high-density MPs were simulated considering only the estuarine area, with MPs being released only from São Gonçalo Channel. In these scenarios (3 and 4), in addition to 311 312 advection and dispersion, the biofouling process was also included in the configuration 313 setup, aiming to indicate the trajectory of high-density MPs. High-density MPs, represent those subject to biofouling, included here to better understand how this phenomenon 314 influences MP dynamics in Patos Lagoon estuary. This decision was driven by the 315 recognition that estuarine regions often exhibit more pronounced interactions with marine 316 currents, which can significantly influence the fate and transport of high-density MPs. In 317 318 all scenarios, the horizontal coefficient (K_h) used were set up to 1, a typical value for coastal systems (Jálon-Rojas et al., 2019). For scenarios involving high-density MPs, 319 biofouling rate (DR) was set to 0.0005 g.cm⁻¹.day⁻¹ (Jálon-Rojas et al. 2019). Simulations 320 321 were run at a 5-minute internal timestep (Δt_i), and only the horizontal displacements due to advection and dispersion were taken into account when presenting the results. 322

Table 1: Simulations setup for the scenarios considering low-density MPs in the entire lagoon (1 and 2) and high-density MPs in the estuary only (3 and 4).

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- 329

	Scenarios			
	1	2	3	4
Domain	Lagoon	Lagoon	Estuary	Estuary
Discharge	Low	High	Low	High
Tributaries	Gua/Cam/SãoG	Gua/Cam/SãoG	SãoG	SãoG
MP type / density	PP/ low	PP/low	PP/ high	PP/ high
Number of MPs released from each tributary	1.200/ 400 / 400	1.200/ 400 / 400	400	400
Predominant wind	NE	SW	NE	SW
Biofouling rate	No	No	0.0005	0.0005

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331 **3 RESULTS**

332 **3.1 HYDRODYNAMICS DURING LOW AND HIGH DISCHARGE PERIODS**

The analysis of the hydrodynamic conditions during the simulations serves as the starting point for understanding the MPs dynamics in the system. The hydrodynamics of the lagoon was assessed using information about the winds (Fig. 2A and 2B), freshwater discharge from the tributaries (Fig. 2C and 2D) and resulting current velocity under predominant wind conditions (Fig. 3).

During the low discharge period (Fig. 2C), maximum values $(2.000 \text{ m}^3.\text{s}^{-1})$ occurred in Guaíba river, and smaller contributions from Camaquã river ($600 \text{ m}^3.\text{s}^{-1}$) and São Gonçalo Channel ($500 \text{ m}^3.\text{s}^{-1}$) (Fig. 1). Winds were predominantly from the north quadrant (Fig. 2A). The highest (lower) wind velocities occur from the NE (E) direction, with speeds exceeding 6 m.s⁻¹ (2 m.s⁻¹).

For simulations during the high discharge period (Fig. 2D), the maximum discharge occurred in Guaíba river (8.000 m³.s⁻¹) and remained high during all the simulation. The same occurs for the Camaquã river, which reached 3.000 m³.s⁻¹, but not to São Gonçalo Channel, which maintained low discharges (800 m³.s⁻¹) throughout the simulation (Fig. 2D). Winds during this period presented greater variability (Fig. 2B), with prevailing wind direction (SW-NE) aligns with the main axis of the lagoon and wind speeds of 9 m.s⁻¹ and 7 m.s⁻¹ from the SW and NE directions, respectively (Fig. 2B).



Figure 2: 3-hour wind velocity and direction during low (A) and high discharge (B) simulated periods, where positive (negative) values represent winds from the south (north) quadrant. Discharges of the three main tributaries of Patos Lagoon during low (C) and high (D) discharge.

Figure 3 presents the resulting current velocity fields under the predominant wind 355 356 directions NE (Fig. 3AB) and SW (Fig. 3CD) observed during the low and high discharge simulations. In the main lagoon water body, the lowest velocities occurred during both 357 low and high discharge (Fig. AC). In the estuarine region, NE (SW) winds promoted more 358 359 intense ebb (flood) currents, especially in the access channel, where the highest velocities occurred (Fig. 3BD). The impact of the winds on the currents is evident, along with the 360 361 notably higher current velocities within the estuarine region compared to the northern sector of the lagoon. It is also possible to observe that through this wind pattern, more 362 363 intense currents are observed along the margins of the lagoon body.



Figure 3: Current intensity and direction in Patos Lagoon during Northeast (AB) and Southwest (CD) wind events. Letters B and D represent the estuarine area.

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368 3.2 LOW-DENSITY MICROPLASTICS TRANSPORT IN THE LAGOON 369 (SCENARIOS 1 AND 2)

370 Figure 4 shows the trajectories of the low-density MPs together with their respective velocities (color scale) during the periods of low (Figs. 4A and 4B) and high (Figs. 4C 371 372 and 4D) discharge in the entire lagoon. During the simulated period, it was observed that the 1200 MPs released from Guaíba river remained in the north of the system, with small 373 dispersion towards the central lagoon. This is highlighted by the low MPs velocities, 374 where most of them did not reach 0.2 m.s⁻¹. Notably, only those MPs reaching the lagoon's 375 shores exhibited higher velocities (0.5 m.s⁻¹). Therefore, even with the predominance of 376 northeast winds (Fig. 2A), the combination between the winds and low discharge (Fig. 377

378 2C) were not enough to enable a significant transport of MPs from the Guaíba river379 towards the central region of the lagoon.

380 During the high discharge period (Figs. 4C and 4D), the dynamics of MPs released by the Guaíba river is more intense. In the early stages of the simulation, a predominance of 381 382 westerly winds was observed (Fig. 2B), which, combined with the high discharge 383 provided by Guaíba river (Fig. 2D), imposed a transport of MPs towards the eastern margin of the lagoon in the northern region, reaching velocities of 0.7 m.s⁻¹. However, 384 similarly to the behavior observed in the low discharge period (Figs. 4A and 4B), the 385 386 center of the northern cell continues to exhibit MPs with low velocities (0.1 m.s⁻¹). At the end of the simulated period, the majority of MPs released from the Guaíba river was 387 388 directed southward, reaching the central region of the lagoon, while the others remained 389 trapped due to the low cellular circulation in the northernmost portion of the lagoon.

390 During the low discharge period, the 400 MPs released from Camaquã river were mainly transported along the western margin of the lagoon, where the higher MPs velocities (0.3 391 m.s⁻¹) were observed, reaching the estuarine area (Figs. 4A and 4B). When arriving at the 392 estuary limit, MPs dispersed near the morphological step that separates the main body of 393 394 the lagoon from the estuary, known as Ponta da Feitoria (Fig. 1C), indicating that this 395 formation acted as a barrier to the transport of MPs towards the estuary. Subsequently, 396 with the prevailing north quadrant winds, currents towards the south prevailed and some 397 MPs entered the estuary via the western margin, reaching the coastal region. Due to the low discharge from the Camaquã river during the simulated period (Fig. 2C), wind 398 399 prevailed as the main contributor to the transport of MPs during this period.

400 During high discharge period, however, the majority of MPs released from Camaquã river 401 remained trapped within the recirculation of the lagoon south cell. Furthermore, the 402 western margin near the mouth of Camaquã river was where the MPs exhibited higher 403 velocities, reaching up to 1 m.s⁻¹. In the central part of the south cell MPs particles 404 remained with low velocities (0.1 m.s⁻¹).

405 Considering the 400 MPs released from São Gonçalo Channel during the low discharge 406 event, it is clear that they encountered stronger currents (Fig. 3) and dispersed south 407 reaching the coastal area with high velocities (Figs. 4A and 4B). Due to the low discharge 408 from the São Gonçalo Channel (Fig. 2C) during this period, the transport of MPs occurred 409 based on the influence of the predominant NE wind (Fig. 2A). However, as MPs approach the narrow 20 km long access channel, the funneling effect intensifies the currents and it
is possible to observe higher velocities of MPs near the mouth (0.8 m.s⁻¹) (Fig. 4B). It is
important to observe that some MPs enter the shallow bays of the estuary.

When looking at the high discharge period (Fig. 4BC), MPs released from São Gonçalo 413 414 Channel also reached the coastal area (Fig. 4D). It was also noticeable that almost the entire estuarine area exhibited MPs with high velocities, differing only where they moved 415 416 to shallower areas and towards the margins. These pronounced displacements of MPs 417 corresponded to the stronger currents that occurred in this area (Fig. 3B). Even with the 418 wind varying in direction during the simulated period (Fig. 2B), the predominant direction of the MPs trajectories was NE, promoting ebb flows towards the coastal area. Some MPs 419 420 particles were also observed in the estuary shallow bays.



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Figure 4: Trajectories of low-density MPs during low (A, B) and high (C, D) discharge.
Letters B and D indicate the estuarine area. Color bar indicated microplastics velocities
during trajectories and no-fill makers the release sources.

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426 3.3 HIGH-DENSITY MICROPLASTICS TRANSPORT IN THE ESTUARY 427 (SCENARIOS 3 AND 4)

Considering the presence of biofouling on the MPs, 400 high-density MPs particles were 428 429 released from São Gonçalo Channel during the low discharge period (Fig. 2C). Results 430 indicated that throughout the simulated period, MPs remained restricted to the upper 431 estuarine region, near their release location (Fig. 5A). Although they were in an area of higher currents (Fig. 3A) and predominant NE winds (Fig. 2A), their dispersion was 432 small. Consequently, MPs velocities were also small compared to the other scenarios, 433 reaching 0.1 m.s⁻¹. In the end of simulated period, no MP reached the coastal area or even 434 435 the lower estuarine region.

In the high-discharge period (Fig. 5B) the high-density MPs also remained close to their release area, with small dispersion. Even with a greater prevalence of SW winds (Fig. 2B), which configure flood currents in this region (Fig. 3B), they had little influence on the trajectories. However, unlike what happens during the low discharge period (Fig. 5A), during high discharge, the velocities of the MPs generally increase, with some particles reaching 0.2 m.s⁻¹, while the others remained constant in 0.1 m.s⁻¹. Just like in low discharge, at the end of the simulated period, no MP reached the coastal area.



Figure 5: Trajectories of high-density MPs released from São Gonçalo Channel during
the low (A) and high (B) discharge simulations. Color bar indicated microplastics
velocities during trajectories and no-fill makers the release source.

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448 3.4 HOT SPOTS - POTENCIAL ACCUMULATION AREAS

Figure 6 present the probability of MPs accumulation in Patos Lagoon (hot spots), 449 450 considering both low and high-density MPs simulated during period of low freshwater discharge. Results of the PDF analysis for the low discharge period (Fig. 4A and Fig. 5A), 451 452 indicated higher likelihood of MPs hot spots occurrence in the three regions of the main lagoon (Fig. 6A): in the northern area of the lagoon, in the central cell close to Camaquã 453 454 river mouth and in the morphological step at Ponta da Feitoria (Fig. 1C). In the estuarine area, a large number of hot spots was observed, mainly near the São Gonçalo Channel 455 456 mouth. Furthermore, hot spots were also observed in the lower estuarine region and 457 towards its mouth, especially near Terrapleno Island and around the mouth of Mangueira 458 Bay (Fig. 6B). It was also possible to observe hot spots on both sides of the West Jetty 459 (Fig. 6C).





Figure 6: Density maps from low and high density microplastics simulated during a low
discharge period. Letters B and C represent the zoom in the red square, where letter B
encompasses the region of Terrapleno Island, and letter C covers the area of the West
Jetty. The color bar indicates the normalized probability values from the density maps.

The probability of hot spots of MPs during the high discharge period when considering 465 low (Figs. 4C and 4D) and high-density MPs (Fig. 5B) is presented in Figure 7'. Results 466 indicated hot spots of MPs near the mouth of Guaíba river (Fig. 7A) and in the 467 468 northeastern region of Patos Lagoon (Fig. 7B). Some hot spots were also observed on the west margin of the south cell due to the Camaquã river contribution. The estuarine area 469 470 also has hot spots near the mouth of São Gonçalo Channel (Fig. 7A), similarly to the low discharge period (Fig. 6A). Other estuarine MPs hotspots were observed at the western 471 472 jetty.





Figure 7: Density maps from low and high density microplastics simulated during the
high discharge period. Letters B and C represent the zoom in the red square, where letter
B encompasses the region around the mouth of Casamento lagoon, and letter C covers
the area of the West Jetty. The color bar indicates the normalized probability values from
the density maps.

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480 4. DISCUSSION

481 **4.1 Low-density microplastics transport in the lagoon (scenarios 1 and 2)**

Simulation results considering only low-density microplastics (Fig. 4) demonstrated that 482 the calculated trajectories for MPs were consistent with variations in wind, discharge (Fig. 483 484 2) and current (Fig. 3) regimes calculated by the hydrodynamic model. Both in the central 485 and estuarine regions, MPs exhibited greater dynamics, managing to reach the coastal 486 area. Furthermore, the different distribution of MPs among the northern, central, and 487 southern sectors of the lagoon also proved to be coherent with the recognized regions of 488 higher and lower hydrodynamics in the lagoon (Paim & Möller, 1986; Niencheski et al., 489 1988). These authors observed weaker currents in the northern and southern cells and more intense ones in the central region of the lagoon, corresponding to a node of a 490 491 standing wave established in the main lagoon body due to local wind action (Moller et al., 2001). Additionally, stronger currents are usually observed on the western margin of
the lagoon. This behavior explains the observed transport of MPs and is consistent with
the results presented by Bortolin et al. (2020) regarding the formation of mud depocenters
in the lagoon.

496 In the low discharge simulated period (Fig. 4AB), it can be summarized that MPs released 497 in the northern part of the lagoon showed a tendency to become trapped within the system, 498 while those released in the central region tended to remain primarily in the estuarine area, with some exportation to the coastal region. The transport of the debris in the north was 499 500 aligned with the low hydrodynamics in the northern cell of the system. They exhibited 501 little displacement along the main axis of the lagoon, primarily following the western 502 margin. Thus, during the simulated period of 15 days, none of the MPs particles released 503 in the northern lagoon reached its central region.

504 MPs released in the estuarine region exhibited a tendency for exportation to the adjacent coastal area. This transport was favored by the interaction between the estuary's geometry 505 506 and the ebb and flood flows driven by barotropic pressure gradients established due to 507 river discharge and wind (Moller et al., 2001). This barotropic circulation pattern was 508 also observed by Schicchi et al. (2023), where it was also indicated that the trajectories 509 of MPs located near river mouths responded markedly to continental discharge. However, 510 in our case, considering MPs distribution from Guaíba river this did not happen. In this 511 region, surface circulation was considered low regardless of the wind or discharge. MPs exported to the coastal area through the access channel had a tendency to either remain 512 513 defined or indefinitely in the recirculation region that occurs south of the western jetty. Cohen et al. (2019) indicated that the concentration of low-density MPs varied spatially 514 515 due to surface circulation. Considering these factors, the distribution of MPs relies significantly on the flow field. 516

517 During the high discharge period (Fig. 4CD), the behavior of MPs differed from that 518 observed during low discharge (Fig. 4AB). It was possible to observe greater 519 displacements of MPs released from the Guaíba river towards the main axis of the lagoon. 520 These pronounced displacements were a response to the higher discharge combined with 521 the wind action. MPs with high velocities moved along the edges of the lagoon, both on 522 the eastern and western shores. Those located in the central regions of the lagoon 523 exhibited lower velocities associated with their shorter displacements. This may be 524 attributed to the extent of the lagoon and its cellular morphology, facilitating the entrapment of particles within the circular dynamics of the region. Also, this distinct
behavior can be explained by the hydrodynamics of the lagoon itself. The higher
velocities along the margins create a pathway for the MPs in response to the wind action.
However, as they move towards the central areas of the lagoon, the transport of debris
slows down with velocity.

However, in the estuarine region, favored by the narrowing of the access channel towards the estuary mouth, the behavior of MPs responds to the variability of the system. Microplastics are exported more intensely and rapidly towards the coastal region during these periods of low hydrodynamic activity. Summers et al. (2023) indicated that there is an increase in MPs accumulation within the bay during low hydrodynamic periods, whereas the opposite effect expels debris from the area, which corroborates with the results indicated in our region.

4.2 High-density microplastics transport in the estuary (scenarios 3 and 4)

538 Considering biofouling-affected microplastics (high density), in both low and high discharges, MPs released from São Gonçalo Channel exhibited a markedly different 539 540 behavior from those free from biofouling action. The difference was evident for low-541 density MPs in the estuarine region (Fig. 4BD) compared to results from biofouling-542 affected MPs (Fig. 5 AB). Regardless of the wind direction and discharge, MPs remained 543 confined within the upper estuary, exhibiting minimal displacement and velocities. When 544 comparing biofouling-free MPs to those affected by this process, it becomes evident that 545 this process strongly influences the dynamics of MPs. Summers et al. (2023) suggested 546 that MPs with lower density had more tendency to be exported to the continental shelf, while higher-density MPs took significantly longer to be exported, remaining trapped 547 within estuaries for extended periods. The authors findings also corroborate with our 548 results when we take into count the trajectories of MPs from all 4 scenarios. Furthermore, 549 by concentrating our efforts on the estuarine zone, we aimed to capture the critical 550 551 interface between the lagoon and the adjacent oceanic environment, where rapid 552 exchanges and mixing occur. Additionally, it was expected that being considered denser 553 particles, they would have less spatial displacement, thus having a greater tendency to 554 remain within the lagoon body.

555 **4.3 Limitations for the study and future work**

Despite the lack of specific parameterizations for the study region, the presented results 556 557 serve as initial indicators of how MPs transportation occurs in Patos Lagoon. Low-density 558 MPs under the influence of a less dynamic system (low discharge, scenario 1) exhibit a 559 trapping gradient from the Guaíba river towards the São Gonçalo Channel. However, when the dynamics of the lagoon become more pronounced in terms of both discharge 560 561 and wind, MPs show greater displacements from the northernmost region to the mouth of the lagoon throughout the estuary. The variation in MPs concentration, starting from 562 563 Guaíba river and extending to the open ocean through the estuary, highlights the 564 significant function of Patos Lagoon as a buffer system before MPs are transported from 565 land to sea (Meijer et al., 2021). In regard to dense MPs, the plastisphere played a 566 significantly more prominent role than the influence of wind and discharge. It is plausible 567 to infer that Rdisplayed a tendency to remain trapped near their release region. 568 Consequently, considering the presence of the plastisphere suggests that the estuary could 569 potentially act as a short-term sink rather than a source of microplastics for the adjacent 570 coastal region, as shown in Figure 5. Despite this retention gradient, upon analyzing the 571 density maps, it was possible to observe that the estuarine region is the area with the 572 highest probability of occurrence of these particles.

It is crucial to emphasize that the results presented in this study stem from the chosen 573 configurations for the simulations, representing a significant step towards the 574 575 understanding of the dynamics of MPs and the contribution of the plastisphere effect in 576 this process. The investigation of the plastisphere is innovative and provides a strong 577 foundation for future research in this area. Model simulations still face some limitations 578 regarding the parametrization for dispersions coefficients, computational costs, and 579 model validation. However, future research should focus on these limitations and other 580 simulations should be conducted using increasingly realistic parameterizations 581 (horizontal and vertical dispersion coefficients), incorporating other processes 582 (deposition, bottom drag, resuspension, and degradation (Jalón-Rojas et al., 2019), and 583 different characteristics for MPs particles (Santos et al., 2023; Alves et al., 2023). 584 Furthermore, it is worth noting that the transport of microplastics may exhibit different 585 behavior when considering the positive and negative phases of the ENSO cycle. Higher 586 discharges are expected during El Niño events, suggesting a higher tendency for 587 microplastics to be exported to the coastal region during this period. Conversely, during 588 La Niña years, the opposite process might occur.

590 **5. CONCLUSIONS**

591 Low-density MPs exhibited behavior consistent with the hydrodynamics of their respective release regions and dynamic configuration of the lagoon. During low 592 593 discharge, short-term simulations revealed a gradient of MP retention within the system, with a larger quantity of MPs being retained in the north under the influence of lower 594 595 discharges and moderate winds. During this period, the estuarine region exhibits the highest dynamics, influencing the increased export of debris to the coastal area. In high 596 597 discharge periods, MPs exhibited a different dynamic, responding to the configuration of 598 the lagoon. Increased wind activity and continental discharges influenced the larger 599 displacement of the MPs. Unlike in low-discharge, they exhibit higher velocities along 600 the shores of the lagoon body. MPs subjected to biofouling action and released in the 601 estuary exhibited distinct behavior from biofouling-free MPs, remaining retained in the upper estuary in both low and discharge simulations. 602

603 Consequently, it is possible to infer that the regions further from the estuary's mouth may 604 be responsible for retaining MPs within the lagoon body. Conversely, the estuarine region 605 tends to facilitate the exportation of low-density MPs to the coastal region through 606 prevailing ebb currents and trap high-density MPs near release locations. However, this 607 area is also impacted by the presence of these debris because everything that enters the 608 lagoon through rivers or poor waste management eventually reaches the estuarine region 609 and remains there for a certain period.

610 The present study represents a significant advance in assessing the transport patterns of 611 microplastics (MPs) in Patos Lagoon. While acknowledging that the particle model 612 employed has to be further validated, it is crucial to highlight that achieving such a task 613 was not possible due to the high-water level observed in the system during the recent flood events (September 2023 and May 2024). Nevertheless, the findings presented are 614 615 pivotal in understanding how MPs are transported from different release locations within the system, aligning with the hydrodynamic behavior. These insights contribute to a 616 617 comprehension of the source-to-sink process, enabling a robust understanding of spatial 618 and temporal patterns in the exportation or retention of MPs within the system.

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